

特提斯洋与古亚洲洋协同演化控制下的塔里木盆地油气富集效应

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摘要:显生宙特提斯洋和古亚洲洋2大构造域演化历史对塔里木盆地产生了巨大的影响,其效应包括原型盆地的形成与沉积充填过程及油气成藏与调整改造等。原特提斯洋和古亚洲洋分别经历了扩张、俯冲-消减和关闭-碰撞造山过程。特提斯域经历了原特提斯、古特提斯和新特提斯3个阶段,具有此消彼长的关系。古亚洲洋在新元古代—古生代多陆-岛-洋的格架下经历了复杂的伸展—聚敛过程。受特提斯洋(昆仑-阿尔金分支洋)和古亚洲洋(南天山洋分支)协同演化控制,塔里木盆地经历了新元古代—早古生代和晚古生代—新生代2个从伸展—聚敛的巨型盆地旋回,与海平面升降和气候旋回匹配,形成了优质烃源岩、多类型规模性储层和盖层,构成了油气形成的优越物质基础。特提斯洋和古亚洲洋多圈层协同演化控制的油气富集效应体现在:①多类型盆地原型与后期构造改造决定了塔里木盆地不同部位油气成藏与富集的差异性;②北部拗陷广泛分布的优质烃源岩,是塔北—顺托—塔中地区大型油气富集区形成的基础;③多期构造运动伴随的构造破裂与古岩溶作用造就了断控缝洞型和岩溶缝洞型2类各具特点且规模发育的储集体;④区内构造与地温场差异演化决定了油气聚集区内油气相态与次生改造的规律性变化。台盆区大型古隆起和不整合面、走滑断裂带、高能相带及其叠加区是油气有利目标区。

关键词:富集效应;成藏要素;变形改造;盆地原型;构造演化;古亚洲洋;特提斯洋

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Hydrocarbon enrichment effects in the non-foreland area of the Tarim Basin under the coevolution control of the Tethys and Paleo-Asian oceans

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Abstract: The Phanerozoic evolutionary processes of the Tethyan and Paleo-Asian tectonic domains have delivered profound impacts on the Tarim Basin, including the formation of prototype basins and sedimentary filling processes, as well as hydrocarbon accumulation and later adjustment and modification. Both the ProtoTethys and Paleo-Asian oceans experienced phases of expansion, subduction, closure, and collisional orogenesis. Specifically, the Tethyan tectonic domain progressed through the Proto-, Paleo-, and Neo-Tethys stages in a row. Meanwhile, the Paleo-Asian Ocean underwent a complex extension-convergence process within a Neoproterozoic-Paleozoic framework involving multiple continents, islands, and oceans. Due to the coevolution of the Tethys Ocean (the Kunlun - Altyn Tagh branch) and the Paleo-Asian Ocean (the South Tianshan Ocean as a branch), the Tarim Basin experienced two extension-convergence megabasin cycles from the Neoproterozoic to the Eopaleozoic and from the Neopaleozoic to the Cenozoic. These cycles, along with eustatic sea-level changes and climatic cycles, facilitated the formation of high-quality source rocks and various large-scale reservoirs and cap rocks, which jointly created a superior material foundation for hydrocarbon generation and accumulation. The hydrocarbon accumulation effects controlled by the coevolution of the Tethys and Paleo-Asian oceans

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are manifested in the following aspects: differential hydrocarbon accumulation and enrichment across different parts of the non-foreland area (i. e. area covered by Paleozoic marine sediments) in the Tarim Basin under a variety of basin prototypes and later tectonic modifications; high-quality source rocks widely seen in the northern depression serving as the foundation for large-scale hydrocarbon enrichment in the northern Tarim, Shuntuoguole low uplift and Tazhong area; two types of large-scale reservoirs with distinct characteristics: fault-controlled fractured-vuggy and fractured-vuggy karst types formed under tectonic fracturing and paleokarstification associated with multi-phase tectonic movements; and the differential evolution of structures and geothermal fields in the accumulation zone dictating the regular changes in hydrocarbon phase state and secondary alterations. Favorable hydrocarbon exploration targets in the non-foreland area include large paleo-uplifts, unconformities, strike-slip fault zones, high-energy facies tracts, and areas where these three factors overlapping.

Key words: enrichment effect, factor controlling hydrocarbon accumulation, deformation and modification, basin prototype, structural evolution, Paleo-Asian Ocean, Tethys Ocean

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现今的塔里木盆地为环形山链所环绕,夹持于南天山、西昆仑和阿尔金造山带之间,北缘为天山弧形山链,南缘为西昆仑-阿尔金弧形山链,地处特提斯构造域和古亚洲洋构造域的交接部位(图1)。塔里木盆地

演化明显受控于新元古代以来与特提斯洋及古亚洲洋相关联的板块构造演化过程。一般认为,紧邻盆地南缘的昆仑山、阿尔金山隶属于特提斯洋构造域,盆地北缘的南天山属于古亚洲洋构造域。自新元古代晚期以

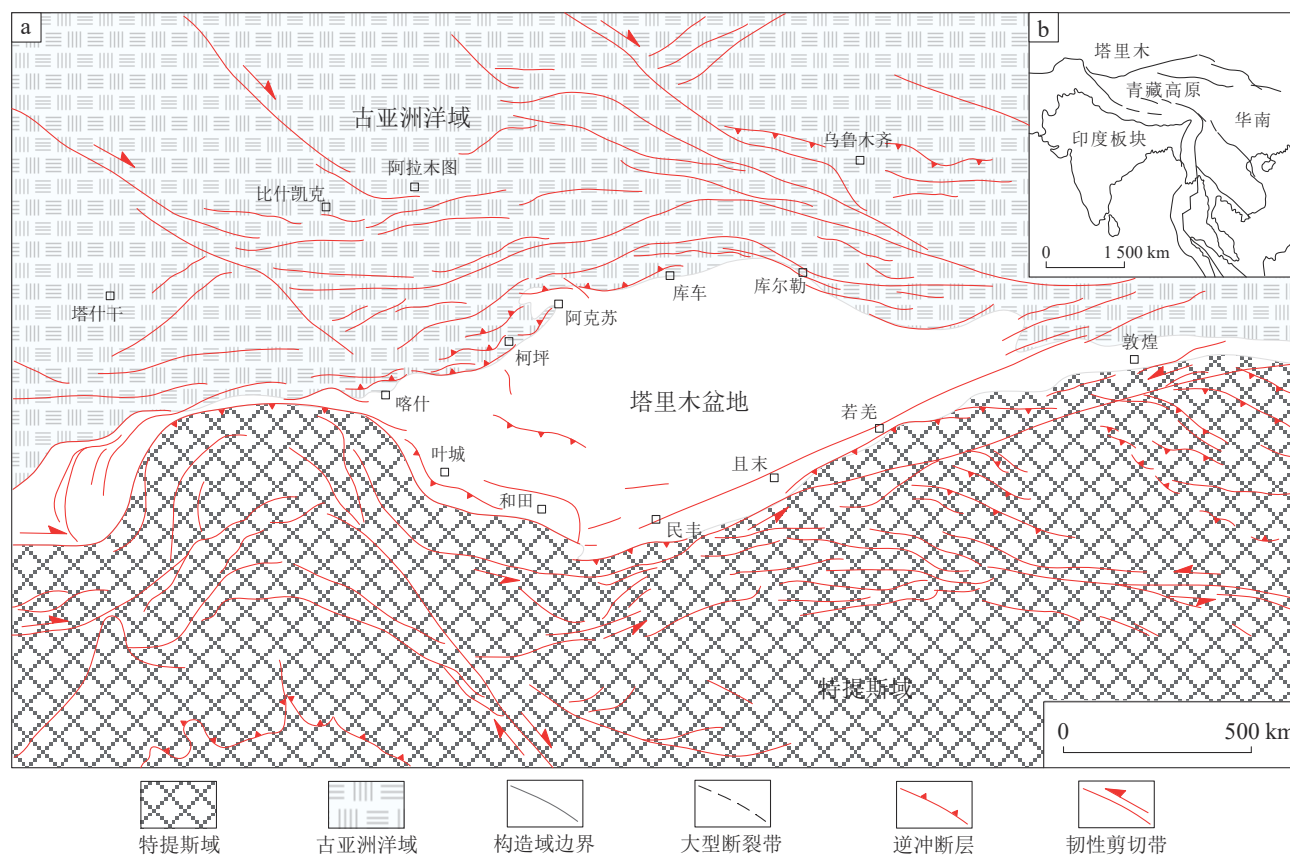


图1 塔里木盆地及周缘地区大地构造单元划分(据参考文献[1]修改)

Fig. 1 Division of geotectonic units in the Tarim Basin and its surrounding areas (modified after reference [1])

a. 亚洲大陆大地构造格架; b. 塔里木邻区构造域划分

来,塔里木地块及周缘地区经历了古亚洲洋盆和特提斯洋盆的开启、俯冲、闭合以及微陆块多次碰撞造山,发生了多期的构造、岩浆及成矿作用。特别是受印度/亚洲碰撞(60.0~50.0 Ma)以来的近程效应和远程效应影响,塔里木盆地周缘发生强烈的隆升、缩短及走滑变形,形成了现今复杂的环形造山系^[1]。研究表明,昆仑山、阿尔金山和南天山自古生代以来,与塔里木盆地长期保持动态的盆-山耦合关系,盆地周缘的洋-陆转换和造山带演化过程主导了塔里木盆地的构造、沉积演化,进而控制了盆地深层海相和中-浅层陆相两套含油气性系统的成藏、富集特征^[1-4]。

塔里木盆地是中国西部典型的叠合盆地之一,含有丰富的油气资源,据估算资源量超过 258×10^8 t油当量(据“十三五”油气资源评价)。历经70余年的艰苦探索,塔里木盆地已发现油气田(藏)70余个,主要分布在库车、塔西南2个新生代前陆盆地和台盆区海相层系之中。其中,环满加尔坳陷西缘分布的沙雅隆起、顺托果勒低隆起和卡塔克隆起是台盆区油气最为富集的地区,先后发现了10余个油气田,探明地质储量超 30×10^8 t油当量。塔河-轮南油田和顺北-富满油气田是这一区域油气田的典型代表(图2)。

塔里木盆地台盆区多年的油气勘探开发成果表明,构造作用形成的断裂体系对该地区的油气成藏与

富集起到极为重要的控制作用。前人对塔北、塔中等地区断裂构造特征进行过大量研究^[5-9],明确了该地区断裂发育具有逆冲挤压和扭动走滑作用的复合效应,断裂形成演化具有多期性、断裂发育具有继承性与区段性的特点。然而,随着勘探的持续深入,更迫切需要进一步梳理构造演化过程及其地球动力学背景,从而更准确地把握该地区的断裂发育规律及断裂控储、控藏作用,为勘探部署提供决策依据。本文从特提斯洋与古亚洲洋协同演化角度出发,集成前期相关研究成果,系统研究特提斯洋和古亚洲洋裂解—聚合过程如何控制塔里木盆地优质烃源岩、多类型规模性储层和盖层发育,阐明油气成藏要素、成藏过程与动力学机制,揭示油气差异富集与特提斯洋和古亚洲洋裂解—聚合的内在联系,旨在为该地区下一步油气的高效勘探提供参考和借鉴。

1 塔里木盆地及邻区特提斯洋与古亚洲洋协同演化过程

塔里木板块北侧边界为汗腾格里-库米什-星星峡缝合线,南侧边界为库地缝合线,东侧边界为车尔臣河-星星峡和阿尔金古转换断层,是一个于新元古代罗迪利亚(Rodinia)大陆解体时形成的独立块体。在

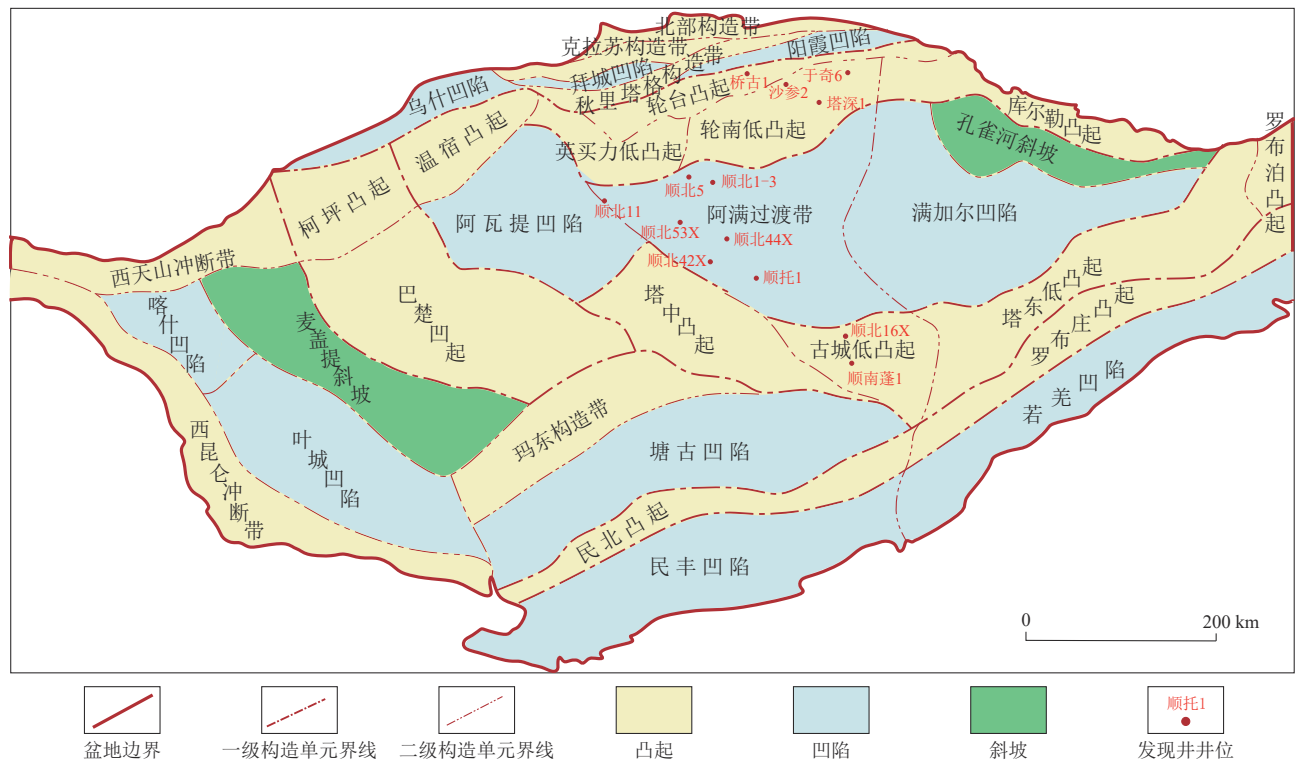


图2 塔里木盆地构造区划与台盆区油气分布

Fig. 2 Structural division of the Tarim Basin and hydrocarbon distribution in the non-foreland area of the basin

显生宙的洋-陆演化历史中,特提斯洋和古亚洲洋2大构造域经历过一系列弧-弧、弧-陆和陆-陆碰撞的造山、造陆作用,塔里木板块实现了与周缘造山带和陆块的复杂拼合过程^[10-12]。这些过程对塔里木盆地内部产生了巨大的影响,其效应包括原型盆地的形成与沉积充填过程、构造多期变形改造及油气成藏与调整改造等。

1.1 特提斯洋(昆仑-阿尔金分支洋)构造演化

1.1.1 原特提斯洋开启(> 540.0 Ma)

原特提斯洋的形成与扩张产生于新元古代罗迪利亚超大陆裂解背景。在阿尔金造山带,已发现索拉克萨依、恰什坎萨依沟等双峰式火山岩,年龄在775.0~750.0 Ma,代表了新元古代大陆裂解^[13-14]。在西昆仑地区,老并一带马尔洋岩组双峰式火山岩组合年龄为603.0 Ma $\pm$ 10.0 Ma^[15],说明震旦纪以前大陆已经裂解。另外,苏巴什附近563.0 Ma $\pm$ 48.0 Ma的洋岛玄武岩^[16]以及阿羌地区545.5 Ma $\pm$ 9.1 Ma的地幔柱型蛇绿岩^[17],指示寒武纪以前原特提斯洋已存在于麻扎-康西瓦-苏巴什断裂一带。阿尔金山目前报道的

最老 MORB 型和 OIB 型蛇绿岩年龄主要在524.0~521.0 Ma,见于安南坝、贝壳滩和阿克塞青崖子等地^[18-19],说明原特提斯洋在早寒武世已进入扩张漂移期。

1.1.2 原特提斯洋俯冲消减(540.0~440.0 Ma)

青藏高原北部地区原特提斯洋的俯冲消减起始于震旦纪末—寒武纪初^[20]。在西昆仑造山带南缘的康西瓦-麻扎缝合带附近,大量岛弧成因的I型花岗岩年龄主要分布在550.0~530.0 Ma^[21-23]。另外,库地、西河休和西奴山等地的一些526.0~519.0 Ma的SSZ型蛇绿岩,指示洋壳向北俯冲导致库地一带弧后洋盆在早寒武世就已经打开^[24-25]。大量构造岩浆事件统计表明(图3b),原特提斯洋的俯冲可能结束于约440.0 Ma的奥陶纪末期,岩石以钙碱性I型花岗岩和岛弧玄武岩为主,幔源岩浆参与贡献大^[17, 26-28]。

在北阿尔金构造带,与早古生代洋壳俯冲有关的SSZ型蛇绿岩和陆缘弧岩浆年龄分别在516.0~448.0 Ma和517.0~449.0 Ma^[18, 29-35],另外,红柳沟一带的榴辉岩和蓝片岩峰期变质年龄512.0~491.0 Ma

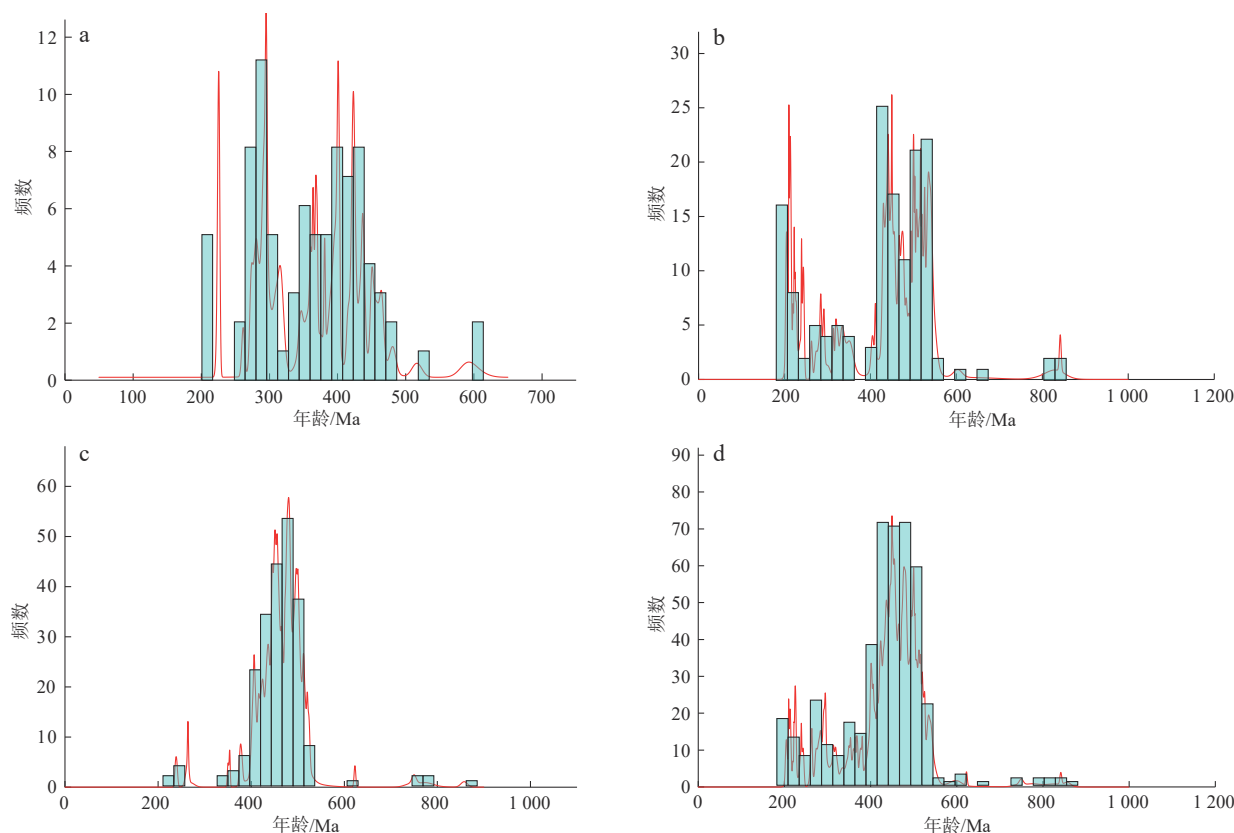


图3 南天山-西昆仑-阿尔金造山带古生代构造岩浆事件年龄谱图

Fig. 3 Age spectra for the Paleozoic tectonomagmatic events in the South Tianshan, West Kunlun, and Altyn Tagh orogenic belts

a. 南天山造山带; b. 西昆仑造山带; c. 阿尔金造山带; d. 南天山-西昆仑-阿尔金造山带

也记录了这一期俯冲事件^[36]。在南阿尔金构造带,早古生代SSZ型蛇绿岩和陆缘弧岩年龄分别在481.0~458.0 Ma和518.0~451.0 Ma^[37-42],而区内一些榴辉岩和麻粒岩等峰期变质年龄也被认为与洋壳俯冲有关,进一步将俯冲事件的峰期限制在约500.0 Ma^[43-46]。晚寒武世500.0~490.0 Ma的构造事件不仅在阿尔金地区表现得最为强烈(图3c),在相邻的祁连山和柴北缘地区也是一次非常强烈的构造运动^[47-49]。

1.1.3 原特提斯洋关闭(440.0~380.0 Ma)

西昆仑造山带一些以S型花岗岩为主的中-酸性岩体,年龄主要在441.0~429.0 Ma^[50-52],指示造山带演化的主碰撞阶段发生在晚奥陶世—早志留世。另外,从塔什库尔干县附近与陆壳深俯冲有关的榴辉岩,以及不孜完沟一带的韧性剪切带中,获得了434.9~428.0 Ma的变质和变形矿物年龄,也印证了主碰撞期主要发生在早-中志留世^[53-54]。直到早泥盆世以后,西昆仑造山带进入了后碰撞演化阶段,库地附近出露的辉绿岩脉和黑云母二长花岗岩被认为是后碰撞伸展环境产物,年龄在408.5~404.3 Ma^[55-56]。

在北阿尔金构造带,主碰撞期和后碰撞期岩浆岩均以I型花岗岩为主,二者年龄分别在453.4~432.4 Ma和434.5~404.7 Ma^[57-62],说明在中志留世就已经处于地壳伸展环境。在南阿尔金构造带,主碰撞期S型花岗岩和后碰撞期A2型花岗岩年龄值分别在462.0~430.4 Ma和435.0~385.2 Ma^[38, 63-69]。另外,还发现一些455.8~445.0 Ma的碱性花岗岩^[38, 70-71],也被认为是后碰撞岩浆产物,说明南阿尔金山在晚奥陶世可能已经处于由主碰撞向后碰撞过渡的演化阶段。

1.1.4 古特提斯洋俯冲消减(380.0~240.0 Ma)

研究表明,龙木措-双湖-怒江-昌宁-孟连对接带是古特提斯主洋盆所在,沿此带大量分布着MORB型、OIB型和SSZ型蛇绿岩岩块^[10]。受古特提斯洋俯冲带后撤控制,西昆仑造山带早石炭世以后进入弧后扩张演化阶段,不仅在奥依塔克、波斯铁列克村和神仙湾等地形成322.8~291.6 Ma的新生镁铁质洋壳^[72-73],还在区域上发育一套以乌鲁阿特组(*C_{1w}*)为代表的双峰式火山岩,代表了弧后裂陷盆地背景^[27]。与南侧俯冲消减有关的晚古生代弧岩浆事件在西昆仑地区也十分强烈(图3b),如库地南、阿卡孜、盖孜和上其汗等地的基性岩脉和中-酸性岩体,年龄主要在337.7~244.1 Ma^[74-78]。

但是在阿尔金造山带还未见晚古生代蛇绿岩和俯冲碰撞有关的变质岩年龄报道,仅在南阿尔金构造带东段茫崖附近的阿克提山发现少量269.2~260.7 Ma的花岗闪长岩,属高钾钙碱性I型花岗岩^[64, 79],说明阿尔金地区石炭纪—二叠纪构造稳定,受古特提斯洋俯冲消减影响并不强烈。

1.1.5 古特提斯洋关闭(240.0~200.0 Ma)

在西昆仑造山带,萨罗依、卧龙岗和泉水沟等地发现的S型花岗岩年龄介于239.5~222.4 Ma^[20, 80-81],以及在班迪尔和塔什库尔干等地获得的242.7~228.4 Ma的榴辉岩变质锆石年龄^[54, 82],都被认为是古特提斯洋闭合后主碰撞期造山产物。另外,区内产出于后碰撞环境的岩浆岩主要分布在大红柳滩、奇台达坂、独尖山和康西瓦铅银矿等地,岩性以二长花岗岩、花岗伟晶岩和双峰式火山岩为主,锆石年龄在227.1~205.0 Ma,说明印支期造山事件直到晚三叠世末期才结束^[83-88]。

阿尔金造山带内与古特提斯洋闭合有关的岩浆事件记录还是很少(图3c),并且之前洋壳俯冲的岩浆活动也不强烈。

1.2 古亚洲洋(南天山洋分支)构造演化

南天山位于汗腾格里-巴伦台-桑树园子-星星峡断裂(即中天山南缘断裂)以南,总体上由古生代海相沉积岩组成,还夹杂一些古生代侵入岩和火山岩,既有代表塔里木北部被动陆缘的海相碳酸盐岩地层,也有代表俯冲增生楔的蛇绿混杂岩沉积,蛇绿岩时代为晚奥陶世—早石炭世,洋盆闭合发生在晚石炭世^[89-92]。

1.2.1 南天山洋扩张(540.0~480.0 Ma)

南天山与塔里木北缘紧邻,在塔里木北缘的阿克苏和库鲁克塔格地区分别发现约755.0 Ma的大陆拉斑玄武岩和755.0 Ma±15.0 Ma双峰式火山岩^[93-94],表明相邻的南天山地区当时也应该处于大陆裂谷环境。受裂谷差异性演化控制,在那拉提南缘断裂附近达鲁巴依和昭苏县夏特南部等地区可能率先拉伸出洋壳,600.0~516.0 Ma的蛇绿岩年龄说明震旦纪早寒武世境外吉尔吉斯斯坦的Terskey洋盆已延伸到伊犁地块南部^[95-96]。

1.2.2 南天山洋早古生代俯冲增生(480.0~380.0 Ma)

中奥陶世以后,南天山洋(或Terskey)洋盆开始俯冲消减^[97]。从构造岩浆热事件记录来看(图3a),南天

山洋向南俯冲并不强烈,为短期、脉冲式或间歇式的正常高角度俯冲,至中泥盆世结束^[98],向北则为长期、多阶段性的俯冲,俯冲过程整体具有自西向东、由早到晚的特点^[99]。作为晚泥盆世以前洋壳俯冲的产物,在科布尔特、哈布腾苏、特克斯科、夏特南部、温泉县城南、色日牙克依拉克、阿拉斯台、巴伦台及乌斯腾高勒河等地发育一些以I型钙碱性为主的弧花岗岩类,部分岩体具有埃达克岩特征^[96],锆石U-Pb年龄主要在481.1~387.0 Ma^[100-105]。与区域上加里东末期的哈萨克斯坦板块拼合事件相对应,在伊犁地块南缘和巴伦台大哈拉军山发生晚志留世—中泥盆世的区域变质作用^[106],以及伊犁—中天山地块的上泥盆统—下石炭统与下伏地层的角度不整合^[107]。

1.2.3 南天山洋晚古生代俯冲增生(380.0~320.0 Ma)

结束早古生代增生造山事件以后,塔里木北缘从晚泥盆世开始进入了一个岩浆活动平静期,沉积了一套上泥盆统—石炭系浅海相地层,鲜有此期间的碎屑锆石年龄或岩体年龄报道,说明塔里木地块北缘已转入被动大陆边缘环境^[108]。而南天山洋向北俯冲仍在继续,只是在晚泥盆世由前进式俯冲转变为后撤式俯冲,相关的弧岩浆岩年龄主要在370.0~317.0 Ma,广泛分布于那拉提山确鹿特达坂、乌孙山、库兹山、巴伦台和辛格尔西等地^[109-110]。与加里东期相比,伊犁—中天山地块南缘海西期弧岩浆岩的中—酸性岩浆岩比例更高,但陆壳物质参与更少,地幔物质贡献增加^[111-112],可能与俯冲后撤引起的弧后伸展有关。在那拉提—巴伦台岛弧岩浆岩带以南,上泥盆统以中—薄层灰岩为主,上部夹石英砂岩,反映了陆棚上的混积碳酸盐岩台地沉积^[113]。至早石炭世末期,南天山洋关闭,两侧的伊犁—中天山地块与塔里木板块发生弧—陆碰撞作用,在和硕县景汗、阿拉套山南缘的温泉地区、木斯陶萨拉南及额尔宾山的盲起苏等地遗留下主碰撞的富铝质S型花岗岩,同位素年龄为310.1~297.2 Ma^[99, 110, 114-115]。

2 大地构造演化控制下的塔里木盆地原型与改造

自新元古代以来,塔里木盆地与古亚洲洋构造体系和特提斯洋构造体系一直处于彼此关联的区域动力学背景下,塔里木盆地与周边造山带的盆地形成、充填过程以及构造变形互为制约、互相印证,成为盆地与周缘造山作用互为响应的直接证据。受特提斯洋和古亚洲洋协同演化控制,塔里木盆地经历了2个巨型的从

伸展—聚敛的盆地旋回(图4),即新元古代—早古生代盆地旋回和晚古生代—新生代盆地旋回。识别盆地中主要的构造不整合,特别是早古生代构造不整合与周缘造山作用的相关关系,对于海相油气资源的开发有着重要的意义。

2.1 新元古代—早古生代盆地旋回

罗迪尼亚超级大陆解体的过程,也是原特提斯洋和古亚洲洋的形成过程,在区域性伸展的宏观大地构造背景下,塔里木板块成为一个被3个新生大洋(南天山洋、昆仑洋和阿尔金洋)所围限的大陆板块,板块边缘形成陆缘裂谷,板块内部则形成了一系列的陆内裂谷,接受了从断陷期碎屑岩到断—拗转化期的碳酸盐岩沉积,构成了盆地内部第一套优质的生—储—盖组合。

T₀地震反射界面是塔里木盆地震旦系与寒武系之间的不整合面,在盆地南部主要表现为寒武系底部玉尔吐斯组或西山布拉克组与前寒武系之间的角度不整合,在盆地北部则为平行不整合接触,对应地质年龄值为约540.0 Ma。康西瓦—麻扎缝合带附近洋盆在550.0~530.0 Ma向南俯冲,以及东昆仑洋约544.0 Ma向北俯冲,被认为是东冈瓦纳大陆北缘的印度—澳大利亚地块群在新元古代—寒武纪巨型陆缘弧(Cadomian弧)活动的产物^[116]。因此,可能是昆仑前锋弧以南的原特提斯洋俯冲主导了震旦纪末期的区域挤压事件,南塔里木广泛抬升、遭受剥蚀,北塔里木发生短暂沉积间断。

在全球缺氧和区域海侵背景下,向盆地腹部超覆沉积了一套含磷、硅质黑色岩系——玉尔吐斯组,该组在盆地北部广泛发育,而在盆地南部分布范围较为有限。之后在总体海退背景下,沉积了肖尔布拉克组白云岩、吾松格尔组膏云岩、沙依里克组灰云岩和阿瓦塔格组膏云岩,纵向上构成了第二个极佳的生—储—盖组合,2套组合在地表露头分别见到了古油藏。

纵观盆地南缘整个早古生代洋—陆演化历史,从寒武纪初期到奥陶纪末期,岛弧、陆缘弧岩浆活动显示持续增强趋势,与此同时,塔里木南缘与昆仑前锋弧之间的昆仑—阿尔金—祁连地区也在经历弧后洋盆的不断扩张与萎缩。在此背景下,塔里木盆地早—中寒武世发育了一个完整的海进—海退沉积旋回,并在晚寒武世又进入另一个海进—海退层序演化阶段。从区域构造岩浆事件年龄来看,二者之间的层序界面可能主要与以阿尔金山地区(~500.0 Ma)为代表的HP-UHP峰期变质事件有关,被认为是原特提斯洋分支洋盆的快速俯冲产物,同期HP-UHP事件还见于祁连山和东

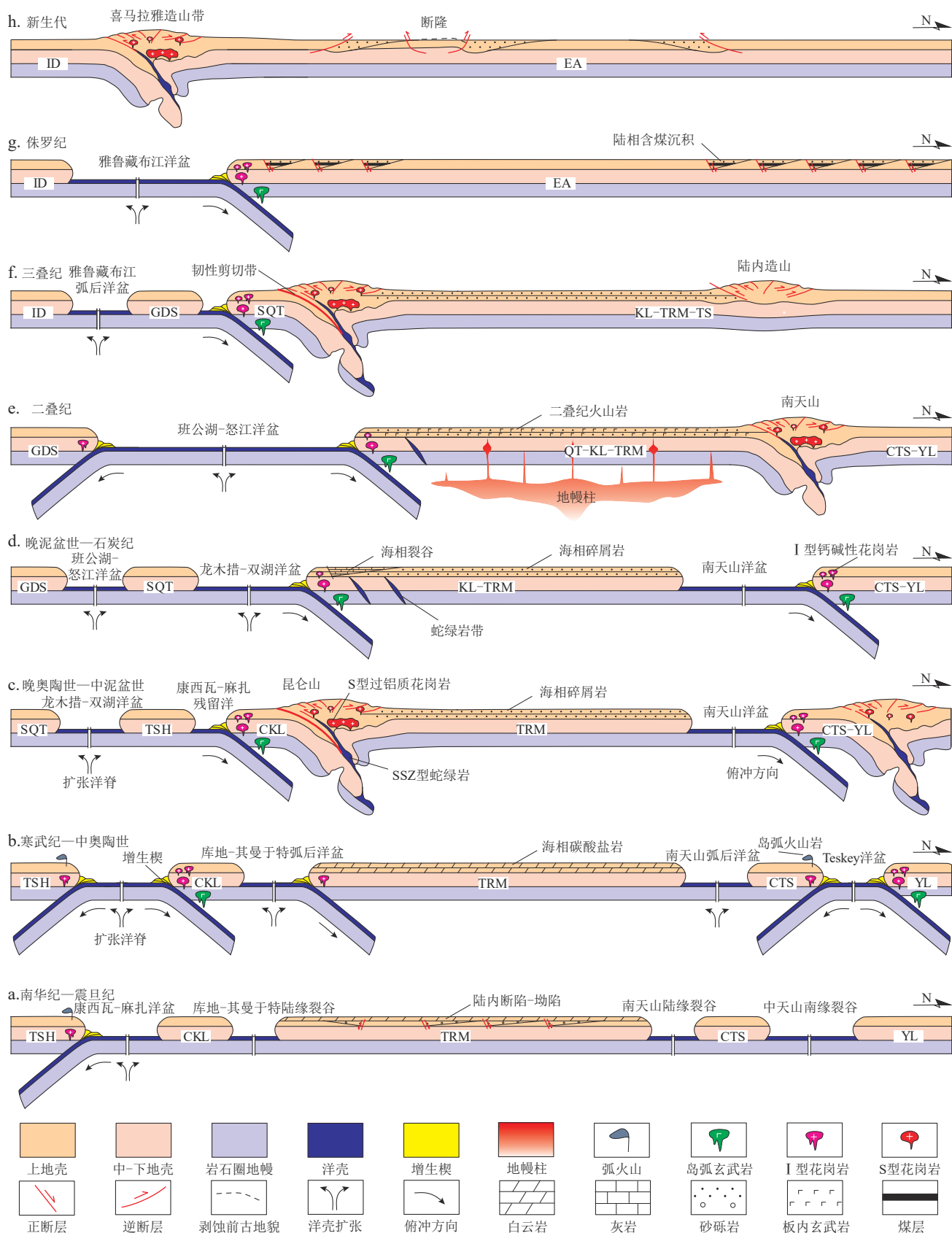


图4 塔里木盆地及周缘地区构造演化模式

Fig. 4 Tectonic evolutionary patterns of the Tarim Basin and its surrounding areas

CKL. 中昆仑地块; CTS. 中天山地块; CTS-YL. 中天山-伊犁地块; EA. 欧亚大陆; GDS. 冈底斯地块; ID. 印度陆块; KL-TRM. 昆仑-塔里木陆块; KL-TRM-TS. 昆仑-塔里木-天山陆块; QT-KL-TRM. 羌塘-昆仑-塔里木陆块; SQT. 南羌塘地块; TRM. 塔里木陆块; TSH. 甜水海地块; YL. 伊犁地块

昆仑造山带。因此,可能是阿尔金及周缘地区寒武纪洋壳俯冲、局部地体拼贴导致了又一次强烈的区域挤压事件,波及整个塔里木盆地,此后在弧后洋盆扩张背景下,盆地内水体变深,受米氏旋回控制,丘里塔格组沉积了上千米厚的碳酸盐岩地层。

T_7^4 地震反射界面是上奥陶统恰尔巴克组泥质灰岩与下伏中奥陶统一间房组生物灰岩之间的盆地级不整合面,对应地质年龄值为约461.0 Ma。中奥陶世,不管是盆地北缘的南天山洋,还是盆地南缘的原特提斯洋,都进入了强烈的俯冲消减阶段,产生了大规模的陆缘弧岩浆活动,并广泛形成一系列弧后洋盆。在昆仑—阿尔金—祁连地区,不仅弧岩浆岩活动在奥陶纪持续增强,还在约478.0 Ma和约461.0 Ma出现2个龄峰值,分别对应 T_7^8 和 T_7^4 地震反射界面,其中, T_7^8 地震反射界面分布在中央隆起带和南塔里木,是中奥陶统鹰山组与下伏中奥陶统蓬莱坝组之间的平行不整合面。在南天山洋,481.0~440.0 Ma的弧岩浆活动也显示逐渐增强的演化特征,直接控制了沙雅隆起中—晚奥陶世的构造抬升和构造不整合面的形成。另外,在祁连、阿尔金造山带(486.3~441.5 Ma)出露以及在南天山地区(450.0~439.3 Ma)分布的SSZ型蛇绿岩,都说明塔里木周缘已进入沟—弧—盆系演化的高潮阶段^[117-118]。因此,盆地边缘幕式俯冲—碰撞作用控制了盆内间歇式构造抬升以及 T_7^8 和 T_7^4 等重要不整合面的形成,并在塘古孜巴斯坳陷、中央坳陷和满加尔坳陷等弧后区形成一个区域性的岩性转换面。塔里木进入周缘前陆盆地与弱挤压背景下的克拉通坳陷演化阶段,南部隆升并发生冲断褶皱变形,东南侧发育大型水下扇。在塔里木东部沉积了数千米的复理石建造。

T_9^0 地震反射界面是志留系底部柯坪塔格组碎屑岩与下伏奥陶系间的盆地级不整合面,对应地质年龄值为约444.0 Ma。奥陶纪末期,随着洋壳俯冲持续进行,昆仑—阿尔金—祁连造山系大部分弧后洋盆和分支洋盆基本都已经关闭,取而代之的是广泛分布的S型花岗岩,此时也是原特提斯构造域岩浆活动最活跃的时期。与此同时,盆地北缘的南天山洋也处于俯冲和弧陆碰撞的过渡阶段,构造活动对整个塔北地区影响显著。在这种南北造山、双向挤压背景下,塔里木陆块南、北两端的和田河古隆起和沙雅隆起相继形成,奥陶系碳酸盐岩遭受剥蚀,发育岩溶不整合和风化壳储层。由于昆仑—阿尔金—祁连地区沟—弧—盆系产物大规模拼贴到塔里木南缘,对整个盆地南部产生了巨大的地壳挤压,包括塔东南坳陷、塘古孜巴斯坳陷和卡塔克隆起,与和田河古隆起连成一片,形成奥陶系与志留系之

间巨大的角度不整合。塔里木盆地北部成为一个受四周新生山系所围限的海湾环境,分别沉积了河流三角洲和潮汐滨岸沉积体系。至此,从新元古代—早古生代分别由陆内、陆缘裂谷—克拉通坳陷、被动大陆边缘—前陆盆地组成的塔里木第一个巨型盆地旋回形成。

2.2 晚古生代—新生代巨型盆地旋回

T_6^0 地震反射界面是上泥盆统东河塘组与下伏不同年代地层之间的又一个盆地级不整合面,对应地质年龄值为约380.0 Ma。经历了志留纪—中泥盆世原特提斯洋和古亚洲洋的区域性俯冲—碰撞造山事件后,南缘古特提斯洋开始扩张,在区域海侵背景下,海水自塔西南进入,向盆地内古隆起超覆沉积,于阿瓦提坳陷、沙雅隆起、卡塔克隆起、塘古孜巴斯坳陷及叶城—和田等地区形成不整合。在班公湖—怒江结合带发现的蛇绿岩(~367.0 Ma),是晚泥盆世古特提斯洋扩张证据^[119]。因此, T_6^0 界面是一个巨大的超覆型不整合,是古特提斯洋扩张背景在塔里木盆地内的产物,也是寻找大规模地层超覆型油气藏的有利层系。

晚泥盆世—中二叠世,塔里木盆地北部—南天山地区,可能存在裂谷环境,盆地内部为弱伸展体制下的克拉通坳陷。中二叠世地幔柱的形成,中—基性岩浆岩的侵入和喷发,对塔里木盆地异常热体制的形成以及大规模的油气生成起到了促进作用。

晚二叠世后,塔里木盆地全面进入陆相沉积演化阶段,南天山前和昆仑山前分别形成周缘前陆盆地和弧后前陆盆地。下三叠统在全盆大部分地区都有分布,但是受南缘古特提斯消亡造山事件的远程影响,中—上三叠统广泛遭受剥蚀,靠近盆地南缘,剥蚀量更大。 T_9^0 地震反射界面是侏罗系与三叠系之间的一个盆地级不整合面,对应地质年龄值为约200.0 Ma。由于盆地北缘南天山在三叠纪隆升,强烈挤压导致了库车前陆盆地的形成,库车坳陷内三叠系发育齐全,并与上覆侏罗系含煤地层连续沉积,呈整合接触。

新特提斯洋扩张过程对处于其北部大陆边缘的塔里木盆地产生了重要影响。进入侏罗纪后,构造—沉积格局发生巨变,造山后的地壳伸展主导了陆内沉降,盆地性质为断陷—坳陷,沉积范围收缩到库车坳陷、满加尔坳陷、塔东南坳陷和塔西南边缘等盆—山结合部位。白垩纪后,塔里木盆地沉积范围进一步扩大,成为一个统一的整体,分别由南、北前陆盆地和中部克拉通坳陷组成,晚白垩世—古近纪,全球性海平面上升使毗邻新特提斯洋的塔西南地区接受海侵。四周山系的不

断隆升,提供了丰富的物源,导致塔里木较大规模的沉降,海水间歇性的侵入和干旱气候背景,沉积了红色碎屑岩与膏盐岩建造组合,是塔里木盆地最重要的区域盖层,对塔里木盆地资源形成、类型与分布产生了关键影响。

新特提斯洋的关闭与青藏高原的形成,导致了中国西部及邻区挤压型盆-山体系的形成,受周缘造山带强烈褶皱隆升的影响,盆地边缘和内部发生了复杂的板缘、板内变形,分别有冲断推覆、滑脱褶皱、走滑压隆-拉分、差异隆升和差异沉降等变形改造方式,控制了油气的形成、调整、改造和破坏过程。

3 特提斯洋与古亚洲洋演化联合控制下的塔里木台盆区油气成藏效应

随着塔里木盆地油气勘探不断向纵深拓展,继塔河-轮南油田主体区之后,相继在托甫台地区、艾丁地区、塔河南部以及哈拉哈塘等外围和深层碳酸盐岩领域不断突破,特别是近几年在多条小滑移距的走滑断裂带内发现了顺北-富满超深层断控型碳酸盐岩油田,证实沙雅隆起-顺托果勒低隆奥陶系整体含油。

塔里木台盆区大规模的油气成藏是特提斯洋与古亚洲洋演化联合控制的结果。这些油气藏均分布在原特提斯演化阶段裂谷和克拉通拗陷分布区,在特殊的古气候和海平面变化背景下,早期发育了优质烃源岩、巨厚的碳酸盐岩地层和大范围展布的膏盐岩盖层^[120],在原特提斯关闭过程中,古隆起和多组断裂带的形成,造就了区域性改造型优质储层的发育展布,为早期大规模油藏的形成奠定了基础。晚古生代-新生代巨型旋回的差异性叠加改造与不同的热体制,是油气性质出现显著差异性又表现出有序性的根本原因。

3.1 塔里木台盆区油气分布

塔河-顺北大型油气聚集区油气相态多样,重质油藏、中质油藏、轻质油藏、挥发性油藏、凝析气藏和干气藏均有钻揭^[121]。超重质油藏的覆盖范围较小,主要分布在塔河油田西北部;重质油藏及中质油藏分布在塔河主体区块;轻质油藏覆盖范围相对较广,主要分布在塔河油田西南斜坡带;挥发性油藏及凝析气藏主要分布在阿克库勒凸起东南斜坡带;干气藏分布在顺南区块和古城区块。不同类型的油气藏在平面上具有“东气西油、南气北油”的分布特征(图5)。

塔河-顺北地区不同走滑断裂带之间以及同一条走滑断裂带不同段之间油藏原油物化性质(颜色、密

度、黏度、气/油比、生物标志化合物和成熟度等)和产能均存在明显的差异性,显示出成藏的复杂性^[122-125]。顺北5号带北段原油密度分布在 $0.828 \sim 0.843 \text{ g/cm}^3$,平均值为 0.834 g/cm^3 ,属轻质原油。顺北5号带中段,顺北1号带分支、顺北1号断裂带原油密度分布在 $0.796 \sim 0.803 \text{ g/cm}^3$,平均值为 0.799 g/cm^3 ,属挥发原油。挥发性油藏相对轻质油藏具有流体密度更低、 C_{7+} 烃组分含量高、油藏饱和和压力相对较小、地-饱压差小、脱气/油比明显更高的特征。造成走滑断裂带不同段之间油藏原油物化性质差异的主要原因可能为:① C_{7+} 烃含量决定了油藏的类型和流体的密度, C_{7+} 烃含量更高的挥发性油藏对应的溶解天然气的能力也更强;② 天然气充注量的差异,由于均处于未饱和状态,仍具有容纳天然气的能力,但是因为不同断裂带天然气充注程度差异有明显区别,因此导致了饱和压力、地-饱压差、脱气/油比等参数出现了明显差异。

3.2 塔里木台盆区走滑断裂特征

基于区域构造活动背景研究和三维地震资料的精细解析,明确走滑断裂具有“线性分段、纵向分层”以及加里东中期、加里东晚期-海西早期、海西晚期“多期活动”的特征^[5-8]。基于野外和钻井的实际解剖,提出断裂多期活动中构造破裂作用形成的断裂空腔和裂缝是主要储集空间类型^[126-129]。

构造裂缝充填的方解石脉记录了构造成岩演化过程,对其开展原位定年分析可以指示断裂形成演化和流体活动的地质时间。本次研究通过对塔里木盆地顺北油田走滑断裂带发育的构造成因方解石脉开展微区原位定年分析,结果指示研究区普遍存在3期断裂活动^[130](图6):① 加里东中期($474.0 \sim 444.2 \text{ Ma}$),在东、西部8条断裂带内普遍发育;② 加里东晚期-海西早期($441.6 \sim 403.2 \text{ Ma}$, $374.4 \sim 326.7 \text{ Ma}$),在顺北5号带以东断裂带内发育普遍;③ 以海西晚期为主($295.9 \sim 232.2 \text{ Ma}$),仅在5号断裂带、4号断裂带和顺南地区的少数样品发育。研究区获得的3期构造成因方解石脉形成时间与基于区域构造活动背景分析和三维地震资料解析所获得的走滑断裂活动时间高度吻合,为揭示塔里木盆地周缘洋陆演化对盆内走滑断裂体系控制提供了直接的时间证据。

3.3 塔里木台盆区油气成藏过程与差异富集

在储层成岩矿物演化序列分析的基础上,通过对发育在成岩矿物中的流体包裹体进行系统分析,将流体包裹体与碳酸盐岩矿物激光原位U-Pb定年、盆地模

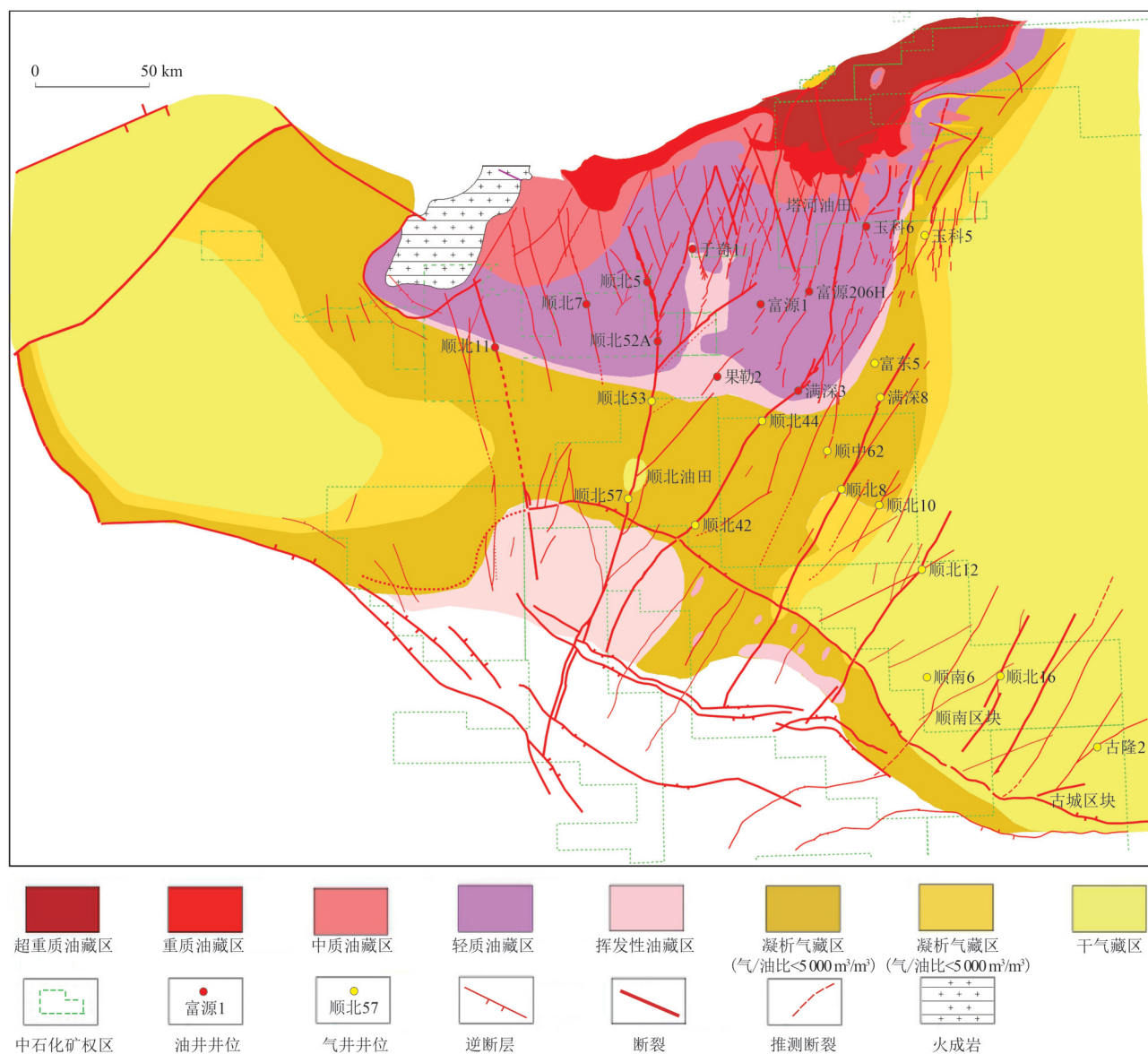


图5 塔里木盆地台盆区油气相态有序分布

Fig. 5 Regular distributions of hydrocarbon phases in the non-foreland area of the Tarim Basin

拟技术、区域构造演化和生烃史相结合,重建了塔河—顺北油气富集区奥陶系深层海相碳酸盐岩油气差异充注演化历史,建立了台盆区深层油气动态成藏模式。研究表明,塔河油田油气藏主要经历了加里东晚期—海西早期、海西晚期、燕山期—喜马拉雅期的动态成藏改造过程(图7)。顺北油气田经历了加里东晚期—海西早期以来的长期成藏过程,在不同构造部位,由于烃源岩热演化和断裂活动性的差异,主要成藏期不尽相同(图7),油藏类型分布也有明显的规律性。

加里东晚期—海西早期,塔里木盆地北部地区抬升暴露剥蚀强烈,导致岩溶缝洞型储集体规模发育,沙雅隆起及邻区的中—下寒武统烃源岩主体处于生油高峰,断裂活动性较强,塔河油田及邻区大规模成藏。同

时,由于储集体埋藏浅,部分地区尚缺乏有效的区域盖层,保存条件差,油藏普遍遭受生物降解的作用。除反映生物降解作用的原油地化特征外,艾丁地区的重质油、钻井岩心在缝洞型储集体中普遍存在的沥青和含沥青的储层流体包裹体,以及缝洞早期充填胶结物中发黄色荧光含烃流体包裹体的较低均一温度,均指示了该期成藏改造过程(图7)。对塔河油田AD19井原油开展有机质Re—Os同位素测年分析,也指示了早期油藏改造发生于 $419.3 \text{ Ma} \pm 6.4 \text{ Ma}$ 。

海西晚期,塔河油田主体区域盖层发育完善,与岩溶缝洞型储集体形成良好的储—盖组合。沙雅古隆起该时期沉降埋藏,斜坡低部位寒武系—中—下奥陶统烃源岩生、排烃,且断裂具有活动性。一方面,奥陶系

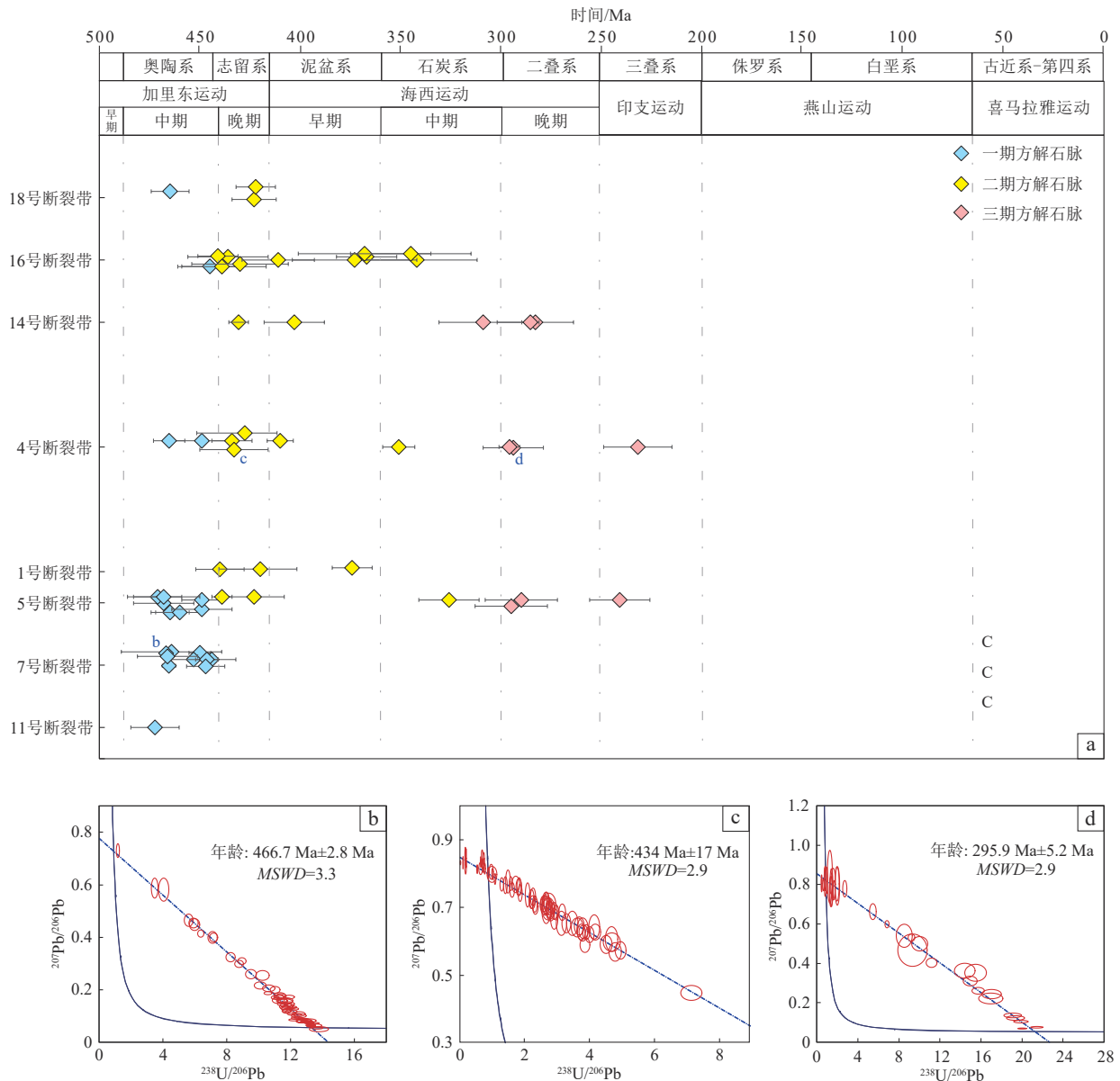


图6 顺托果勒低隆起不同断裂带中-下奥陶统储层方解石脉微区原位U-Pb年龄分布特征和代表性U-Pb年龄
(据参考文献[130]修改)

Fig. 6 Distribution of in-situ calcite veins U-Pb ages in the Middle-Lower Ordovician reservoirs in different fault zones of the Shuntuoguole low uplift and several representative U-Pb ages (modified after reference [130])

a. 不同断裂带方解石脉U-Pb年龄分布特征; b. 7号断裂带第一期方解石脉年龄; c. 4号断裂带第二期方解石脉年龄; d. 4号断裂带第三期方解石脉年龄
(MSWD为平均标准权重偏差。)

岩溶缝洞型油气藏可以再次充注较高成熟度的油气; 另一方面, 早期形成的油气藏可能发生调整, 在上覆层位圈闭中形成新的油气藏。在塔河主体、托普台、跃进和顺北地区都观察发现发黄绿色荧光的含烃流体包裹体, 其伴生的盐水包裹体均一温度指示海西晚期油气充注(图7)。

燕山期—喜马拉雅期, 沙雅古隆起快速沉降埋藏, 斜坡部位下寒武统烃源岩再次生、排高成熟油气, 坳陷

区寒武系一中—下奥陶统烃源岩生、排天然气。该时期, 高成熟油气沿断裂由南部和东部向塔河油田主体部位运聚。塔河东部和于奇东地区表现为晚期气侵特征。在托普台和跃进地区, 大量发蓝绿色荧光的含液态烃包裹体和含气态烃包裹体, 其伴生的盐水包裹体均一温度指示该期成藏改造过程发生于燕山期—喜马拉雅期(图7)。

顺北油气田西部以油藏为主。由于该区始终处于

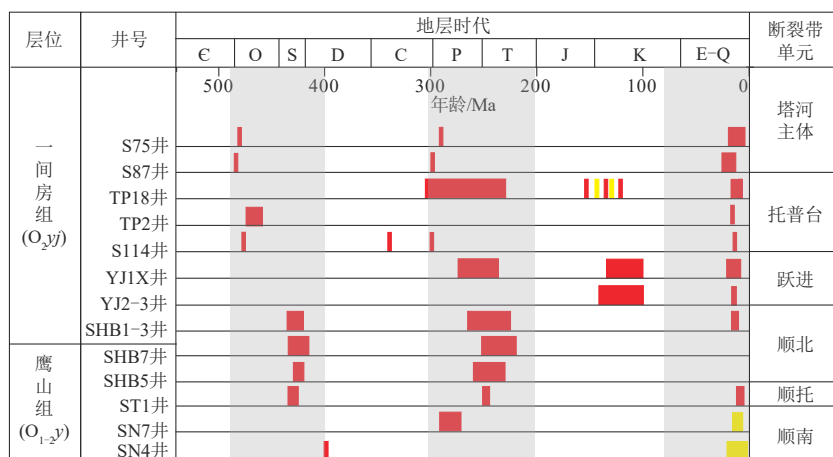


图7 塔河油田—顺北油气田不同地区典型钻井成藏期对比

Fig. 7 Comparison of hydrocarbon accumulation periods in typical wells in different areas of the Tahe oilfield and the Shunbei oil and gas field
(图中红色为油充注,黄色为气充注,灰色区域为不同井对比后确定的主要成藏期。)

构造低部位,且具有持续的低地温场背景^[131-133]。下寒武统烃源岩自加里东晚期—海西早期以来持续生烃。伴随克拉通内小滑移距走滑断裂的发育过程,在寒武系—奥陶系厚层碳酸盐岩中形成了规模性断控储集体,断裂活动促使油气向储集体发生聚集,形成断控缝洞型油气藏。加里东晚期—海西早期油气藏埋藏较浅,断裂活动性强,部分油气藏可能遭受破坏形成储层沥青。此后油藏持续埋藏,次生改造作用较弱,伴随断裂多期活动,不同成熟度的油气可发生多次充注。顺北1号断裂带和顺北5号断裂带北段多口井储层流体包裹体与断裂活动差异的研究指示,加里东晚期—海西早期、海西晚期发生过油气充注;断裂活动强度与油气充注密切相关,在一定程度上控制着油气性质及其分布^[122, 127, 134-135]。

顺南地区以气藏为主。该区具有较高的地温场背景^[124, 131],本区及邻区寒武系烃源岩生烃演化过程较顺北一区早。喜马拉雅期以来以生干气为主,埋深8 000多米的奥陶系碳酸盐岩温度已超过160℃,原油裂解作用发生,早期形成的油藏发生相态转化。顺南1井原油检出较高含量的金刚烷,同时发现含沥青包裹体,与含甲烷包裹体伴生的盐水包裹体均一温度和现今地温基本相当,指示喜马拉雅期天然气成藏。顺北42X井与烃类伴生盐水包裹体均一温度主要分布在90~110℃和130~140℃,指示在该区也存在加里东晚期和海西晚期2期液态烃成藏过程(图7)。

根据塔里木台盆地塔河油田和顺北油气田形成的地质条件和成藏过程,可以看出优越的源—储条件和时—空配置关系是大型油气田形成的关键要素。受塔里木盆地南部原特提斯洋(昆仑—阿尔金分支洋)关闭

影响,盆地内部经历了加里东中、晚期多期挤压走滑变形改造,所形成的“古岩溶地貌和断裂”控制了塔河和顺北2个大油气田油气的局部富集。在储集体与油气藏形成演化的整个过程中,构造作用尤为关键。

综上所述,特提斯洋和古亚洲洋多圈层协同演化控制的塔里木台盆地油气富集效应体现在:①多类型盆地原型的并列叠加与后期构造改造方式和强度,决定了不同构造单元油气成藏要素与过程的差异性;②塔里木北部稳定沉降和高生产力古环境形成了广泛分布的优质烃源岩,是塔北—顺托—塔中地区大型油气富集区形成的基础;③多期构造运动与适度的抬升改造伴随的多期构造破裂与古岩溶作用,造就了2类各具特点且规模发育的缝洞型储集体;④区内构造与地温场差异演化,决定了油气聚集区内油气相态与次生改造的规律性变化。大型古隆起和不整合面、成排的走滑断裂带和高能相带及其叠加区是油气有利目标区。

4 结论

1) 原特提斯洋和古亚洲洋的形成与扩张产生于新元古代罗迪利亚超大陆裂解背景。随后分别经历了扩张、俯冲—消减和关闭—碰撞造山过程,特提斯洋(昆仑—阿尔金分支洋)经历了原特提斯、古特提斯、新特提斯3个阶段,具有此消彼长的关系。古亚洲洋(南天山洋分支)在新元古代—古生代的多陆—岛—洋的格架下经历了复杂的伸展—聚敛过程。2个构造域分别或联合作用,为塔里木板块的构造—沉积作用提供了宏观背景。

2) 受特提斯洋和古亚洲洋协同演化控制,塔里木盆地经历了2个巨型的从伸展—聚敛的盆地旋回,分别为新元古代—早古生代盆地旋回和晚古生代—新生代盆地旋回。其中,由陆内、陆缘裂谷—克拉通拗陷、被动大陆边缘—前陆盆地组成的塔里木第一个巨型盆地旋回,与海平面升降和气候旋回匹配,形成了优质烃源岩、多类型规模性储层和盖层,构成了油气形成优越的物质基础。第二个巨型盆地旋回也分别形成了多个规模不等的含油气系统,同时,后期对前期盆地的叠加改造,为油气调整改造和分布提供了宏观背景。

3) 特提斯洋和古亚洲洋多圈层协同演化控制的油气富集效应体现在:①多类型盆地原型与后期构造改造决定了塔里木盆地不同部位油气成藏与富集的差异性;②北部广泛分布的优质烃源岩,是塔北—顺托—塔中地区大型油气富集区形成的基础;③多期构造运动伴随的构造破裂与古岩溶作用造就了2类各具特点且规模发育的缝洞型储集体;④区内构造与地温场差异演化决定了油气聚集区内油气相态与次生改造的规律性变化。台盆区大型古隆起和不整合面、走滑断裂带、高能相带及其叠加区是油气有利目标区。

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